



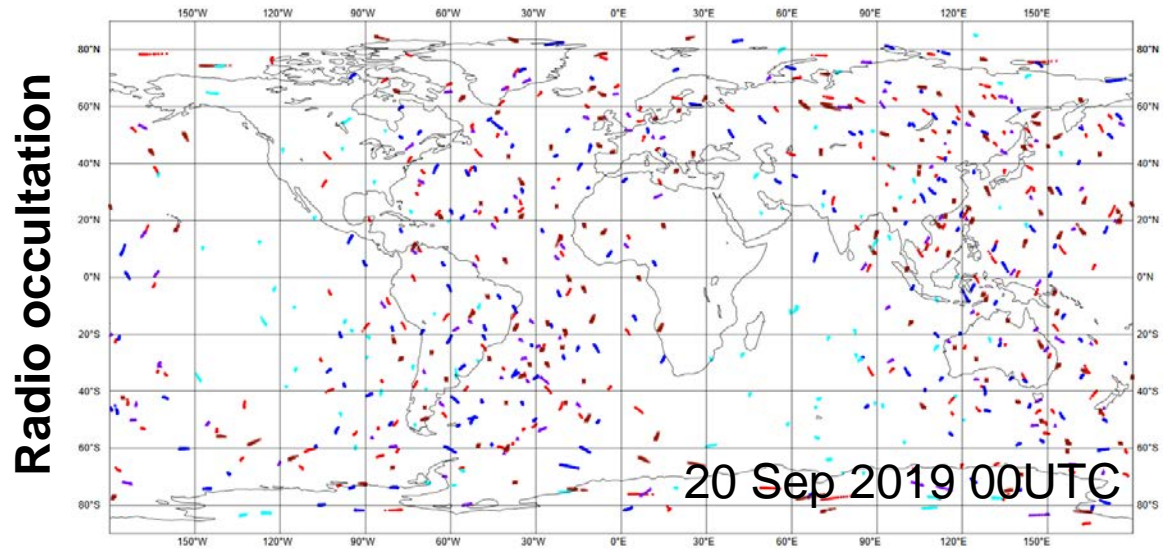
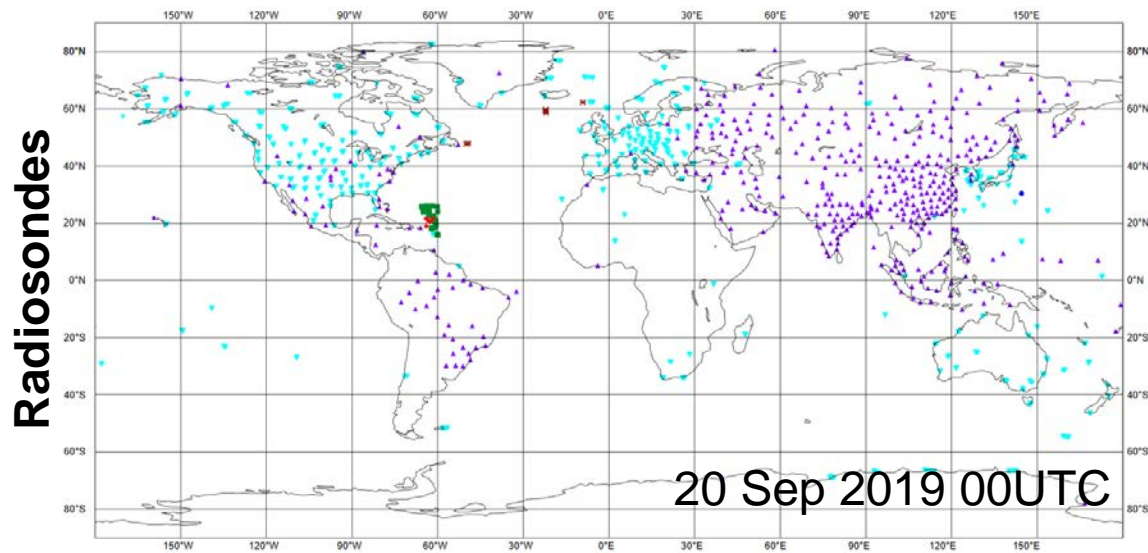
Towards an unbiased stratospheric temperature analysis

Patrick Laloyaux, Massimo Bonavita, Marcin Chrust, Mohamed Dahoui, Jacky Goddard, Selime Gürol, Sean Healy, Elias Holm, Simon Lang, Inna Polichtchouk



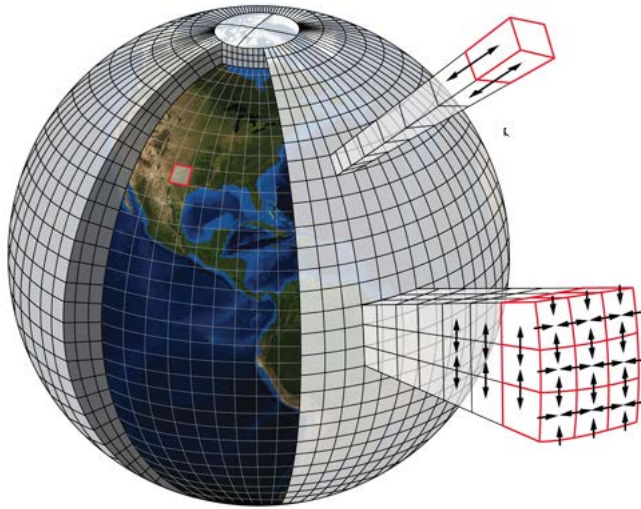
EUROPEAN CENTRE FOR MEDIUM RANGE WEATHER FORECASTS

How we observe stratospheric temperature

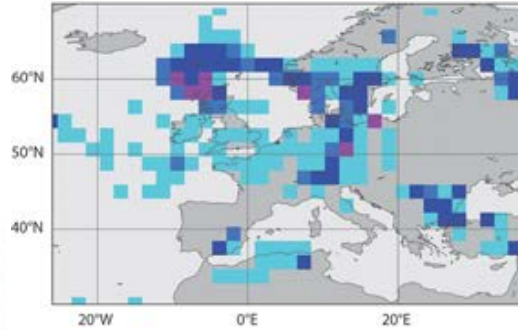


Radiosondes and RO are used as anchoring observations in the atmosphere

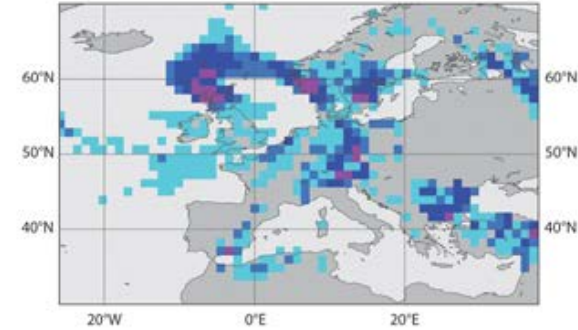
How we model stratospheric temperature (IFS)



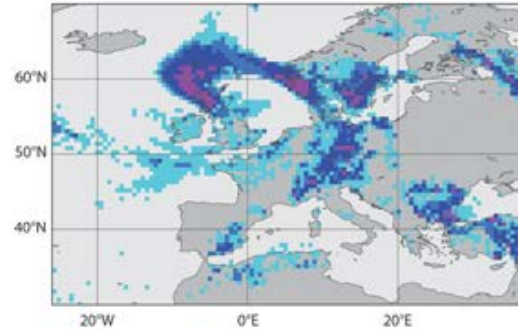
210 km



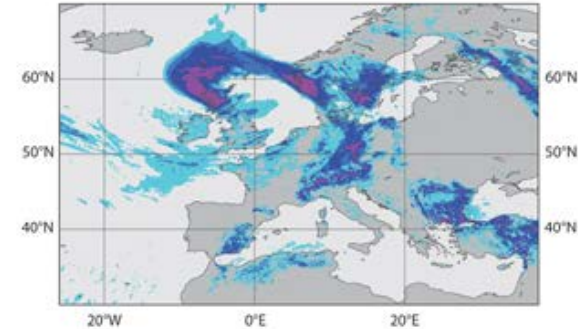
125 km



62 km



9 km / L137 / 0.01h Pa



Model upgrades include

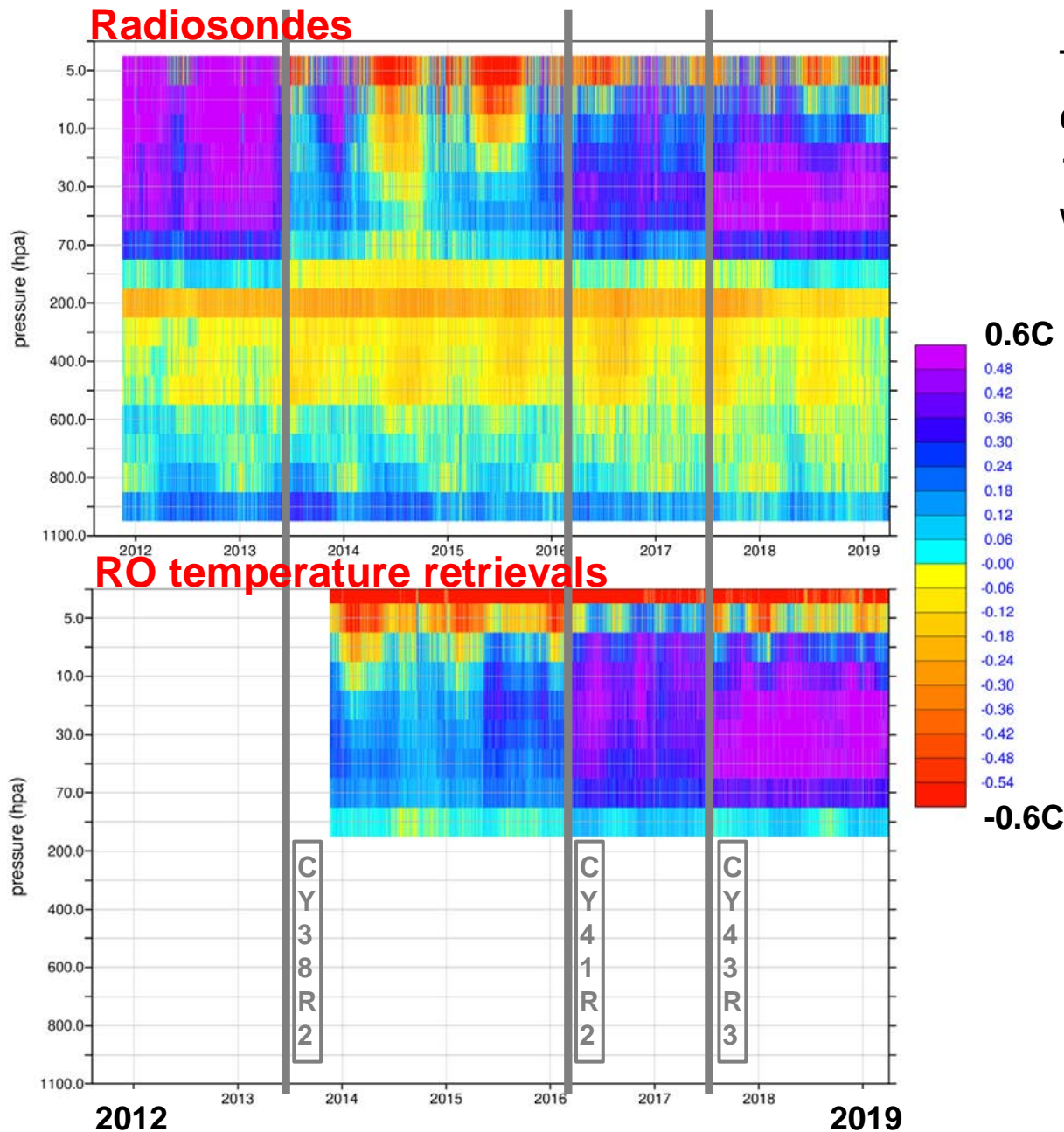
- resolution
- dynamic
- physic

Purpose of this talk is to demonstrate a new application of RO

→ RO used to diagnose temperature model biases

→ RO used to correct automatically those model biases in 4D-Var

Temperature bias in the operational IFS model (1/2)



The short-term model bias is estimated by comparing the 12-hour first-guess trajectory with radiosondes and GPS-RO

Similar signal with the two types of observations:

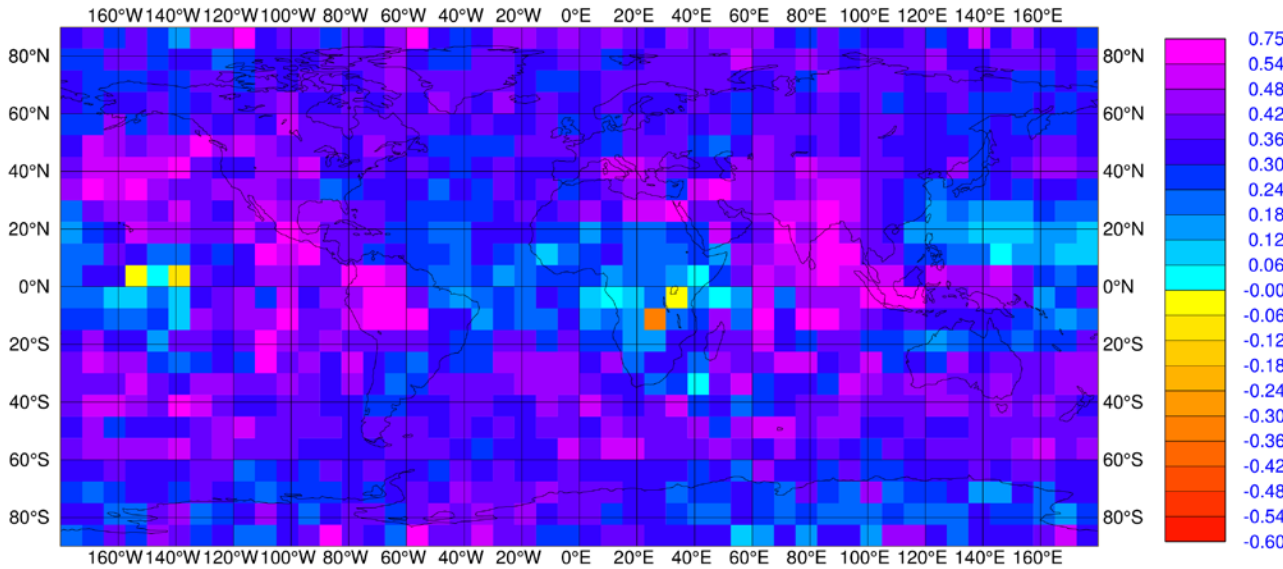
→ bias reduced with new vertical resolution (L137 in CY38R2)

→ bias increased with new horizontal resolution (Tco1279 in CY41R2)

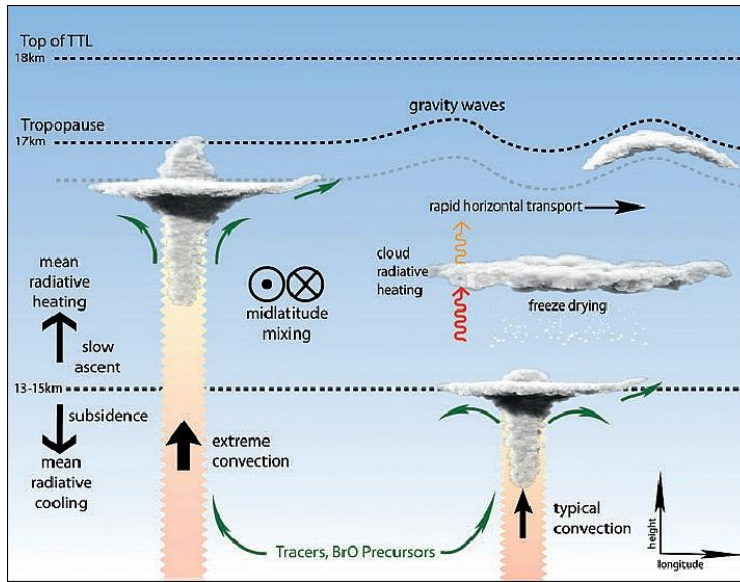
→ bias increased with new radiative scheme (CY43R3)

Temperature bias in the operational IFS model (2/2)

Temperature first-guess departure with respect to RO (~70hPa, January 2017)



→ model error is large scale and presents specific features

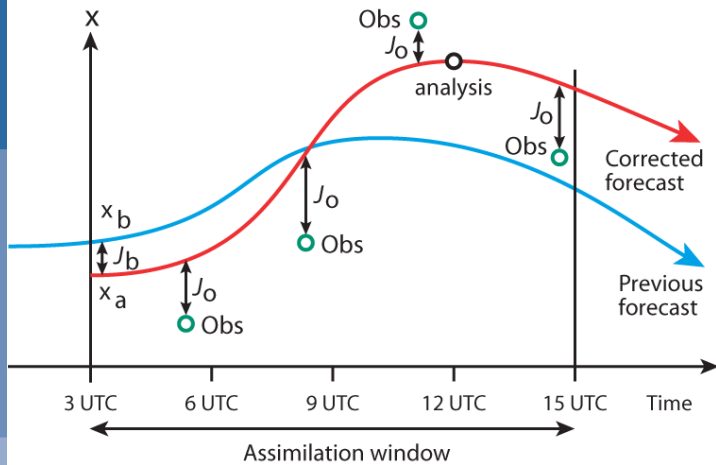


The bias is due to

- discretization errors in the vertical advection (dynamical core)
- inadequate representation of gravity waves in the vertical direction

RO can be used to diagnose the spatial structure of the model error

4D-Var assimilation (1/3)



Strong-constraint

- First-guess trajectory
- Observations
- Compute a correction at initial time
- Analysis trajectory

The model is assumed to be perfect (strong-constraint)

$$\mathbf{x}_k = \mathcal{M}_{k,k-1}(\mathbf{x}_{k-1}) \quad \text{for } k = 1, \dots, N$$

Cost function depends only on the state at the beginning of the assimilation window

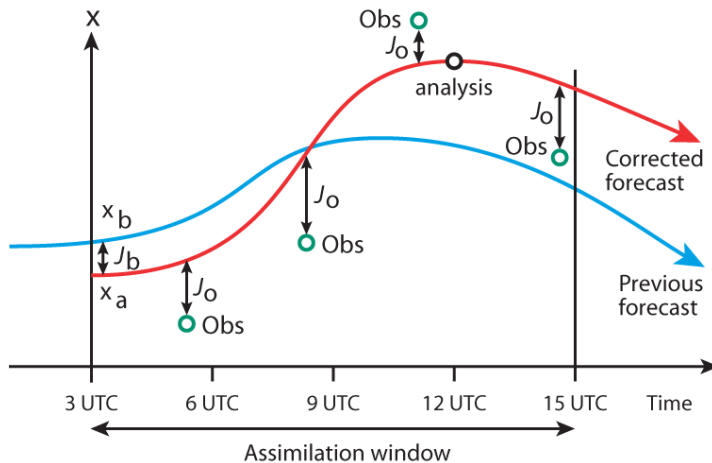
$$J(\mathbf{x}_0) = \frac{1}{2} (\mathbf{x}_0 - \mathbf{x}_0^b)^T \mathbf{B}^{-1} (\mathbf{x}_0 - \mathbf{x}_0^b) + \frac{1}{2} \sum_{k=0}^N (\mathcal{H}_k \mathcal{M}_{k,0}(\mathbf{x}_0) - \mathbf{y}_k)^T \mathbf{R}_k^{-1} (\mathcal{H}_k \mathcal{M}_{k,0}(\mathbf{x}_0) - \mathbf{y}_k)$$

4D-Var assumes **random zero-mean errors** for all sources of information, but the IFS model has biases

4D-Var assimilation (2/3)

Unknown forcing is introduced (additive, Gaussian, constant within the assimilation window, no cross-correlation with the background error).

$$\mathbf{x}_k = \mathcal{M}_{k,k-1}(\mathbf{x}_{k-1}) + \eta \quad \text{for } k = 1, \dots, N.$$



Weak-constraint

- First-guess trajectory
- Observations
- Compute a correction at initial time
- Compute a model forcing η
- Analysis trajectory

Cost function depends on the initial state and the model forcing

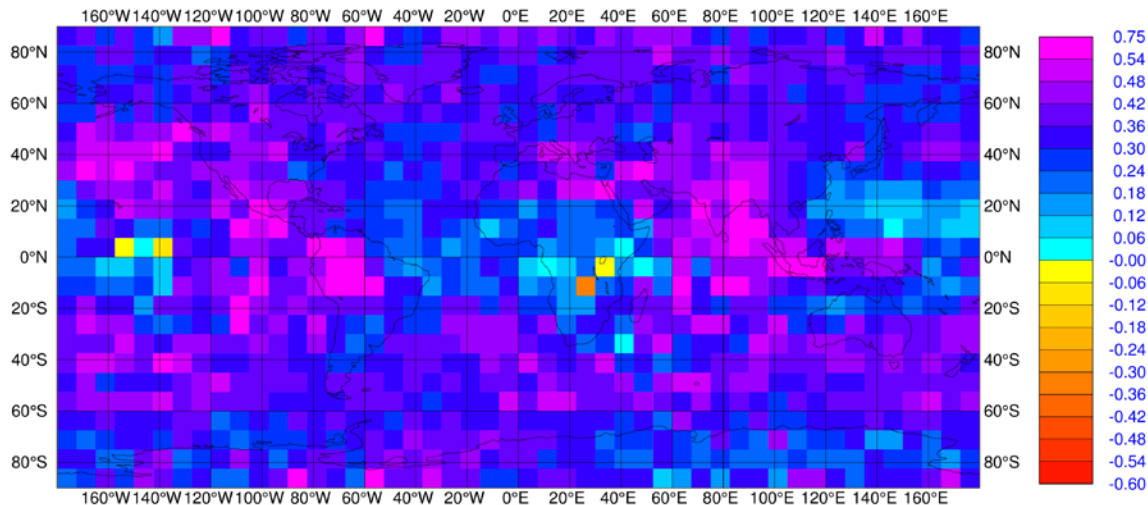
$$\begin{aligned} J(\mathbf{x}_0, \boldsymbol{\eta}) &= \frac{1}{2} (\mathbf{x}_0 - \mathbf{x}_0^b)^T \mathbf{B}^{-1} (\mathbf{x}_0 - \mathbf{x}_0^b) \\ &+ \frac{1}{2} (\boldsymbol{\eta} - \boldsymbol{\eta}^b)^T \mathbf{Q}^{-1} (\boldsymbol{\eta} - \boldsymbol{\eta}^b) \\ &+ \frac{1}{2} \sum_{k=0}^N (\mathcal{H}_k(\mathbf{x}_k) - \mathbf{y}_k)^T \mathbf{R}_k^{-1} (\mathcal{H}_k(\mathbf{x}_k) - \mathbf{y}_k) \end{aligned}$$

4D-Var assimilation (3/3)

$$J(\mathbf{x}_0, \boldsymbol{\eta}) = \frac{1}{2} (\mathbf{x}_0 - \mathbf{x}_0^b)^T \mathbf{B}^{-1} (\mathbf{x}_0 - \mathbf{x}_0^b) + \frac{1}{2} (\boldsymbol{\eta} - \boldsymbol{\eta}^b)^T \mathbf{Q}^{-1} (\boldsymbol{\eta} - \boldsymbol{\eta}^b) + \frac{1}{2} \sum_{k=0}^N (\mathcal{H}_k(\mathbf{x}_k) - \mathbf{y}_k)^T \mathbf{R}_k^{-1} (\mathcal{H}_k(\mathbf{x}_k) - \mathbf{y}_k)$$

→ method described in Sasaki (1970), never worked properly at a NWP centre

→ error statistics (B and Q) need to be specified



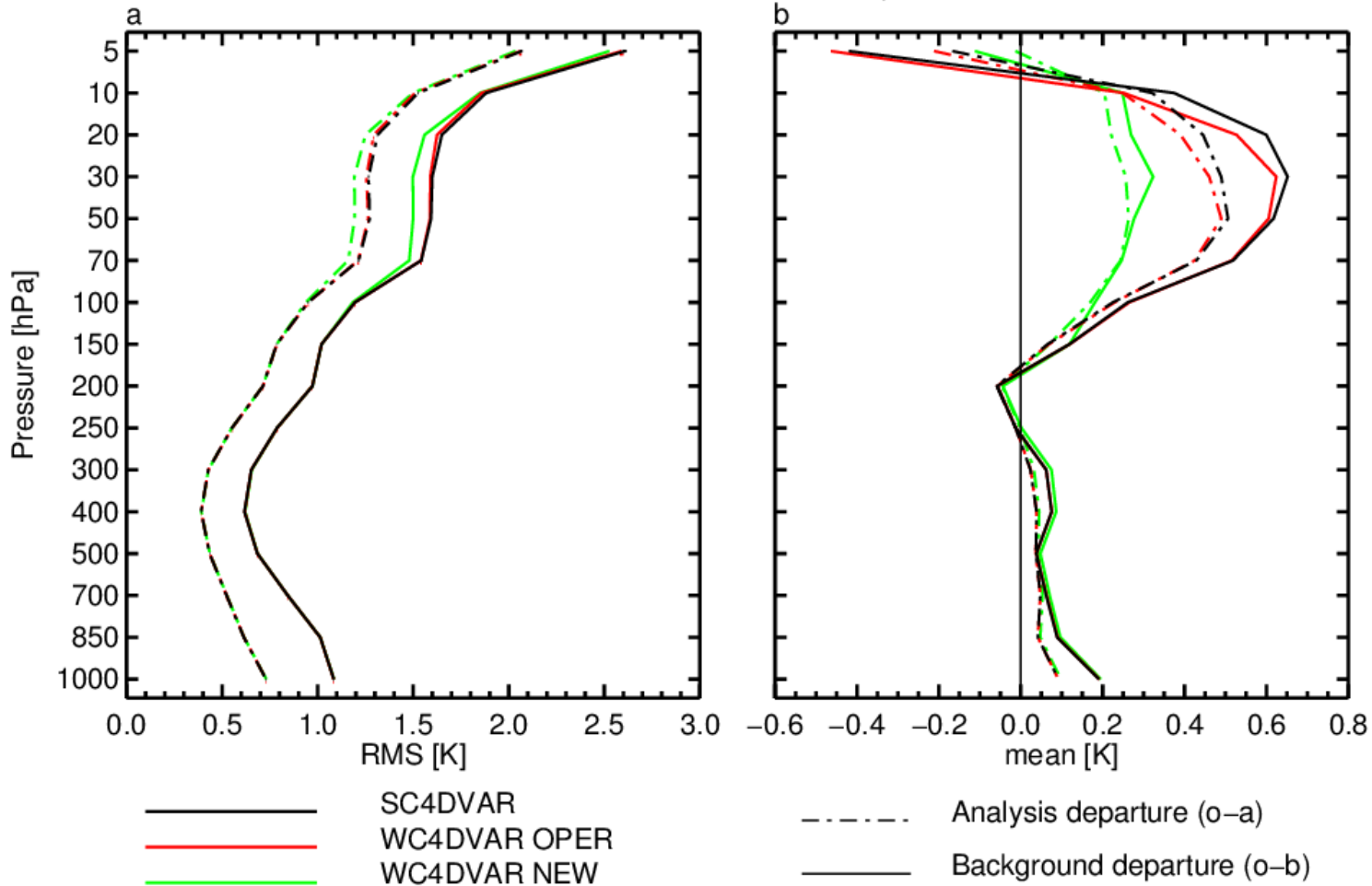
RO shows that model error (Q) is large scale

EDA shows that background error (B) is small scale

→ Scale separation

Weak-constraint 4D-Var with scale separation in IFS (1/2)

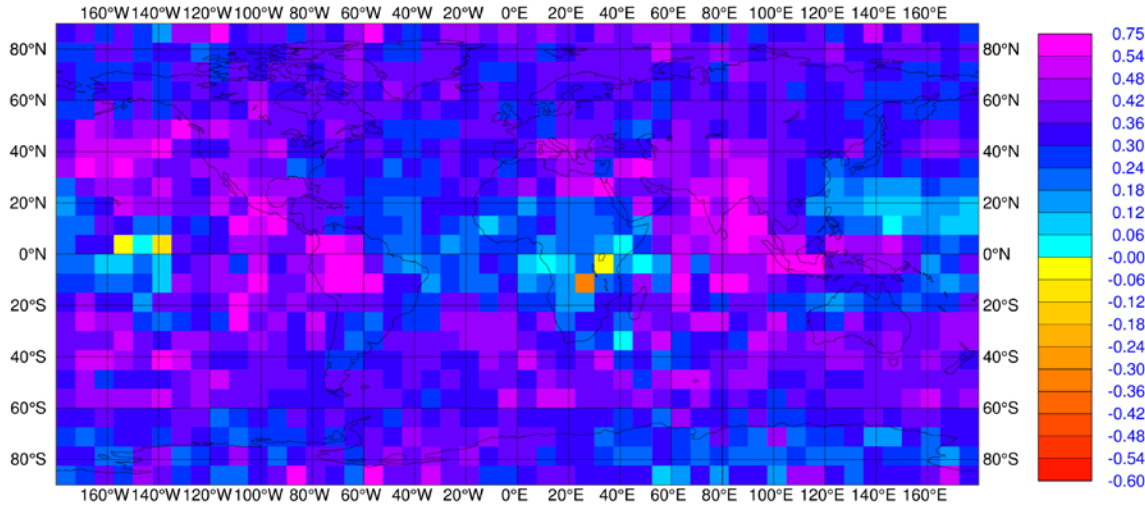
Instrument(s): TEMP-T Area(s): N.Hemis S.Hemis Tropics
From 00Z 1-Oct-2018 to 00Z 14-Apr-2019



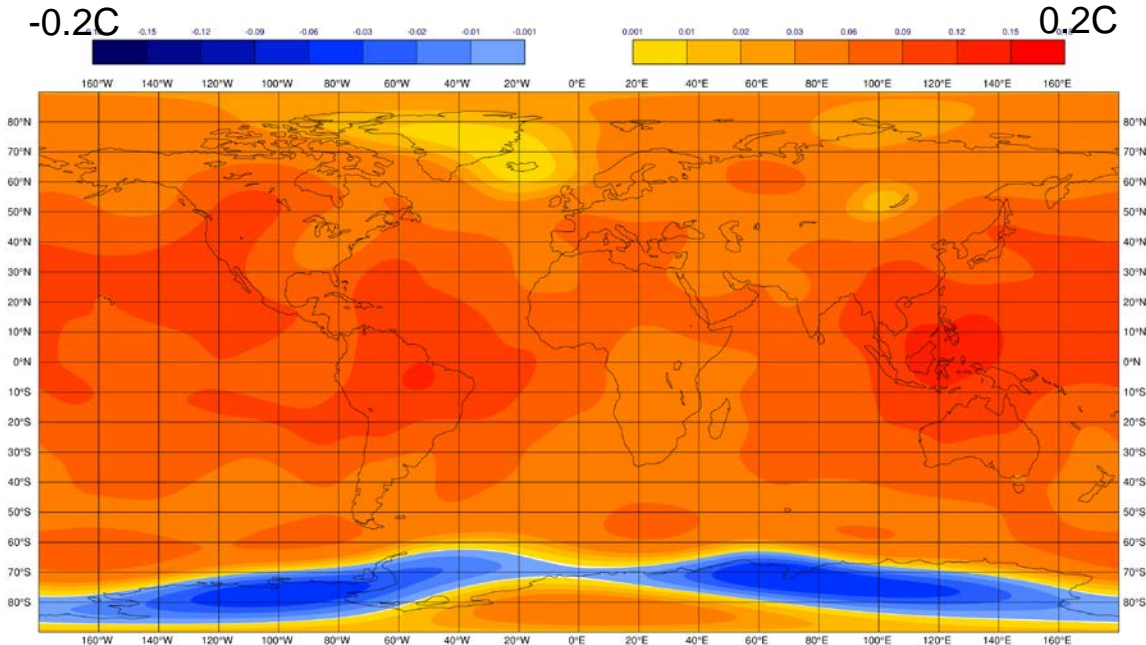
→ temperature bias is reduced up to 50% (0.6C to 0.3C) with respect to radiosondes and GPS-RO

→ temperature RMSE is reduced by 6%

Weak-constraint 4D-Var with scale separation in IFS (1/2)



Cold biases in the lower/middle stratosphere over strong convective regions

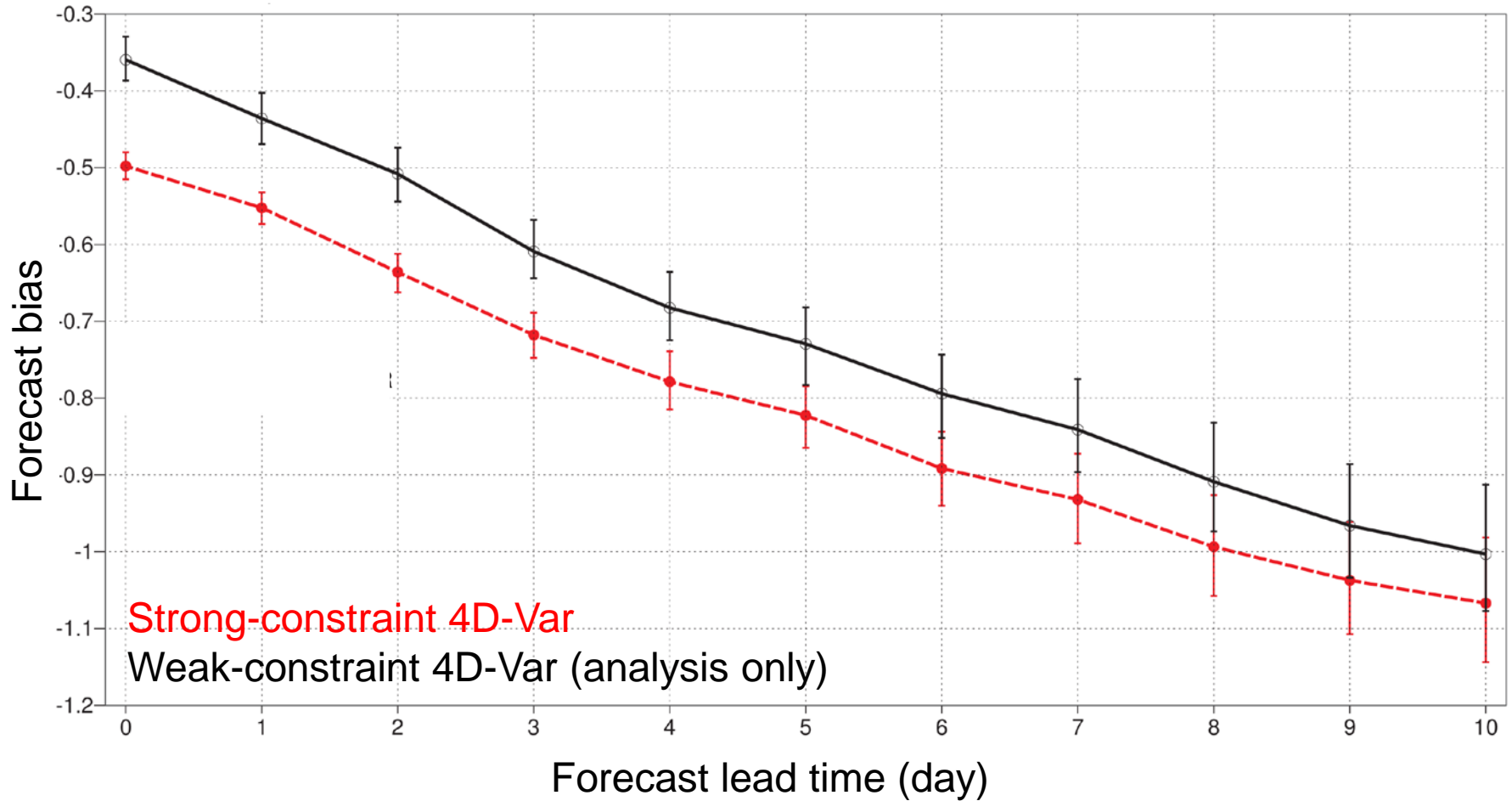


Model error forcing from WC4DVAR at 70 hPa

→ correcting the bias from the missing gravity waves and the dynamical core

Impact of new weak-constraint 4D-Var analysis on forecast skills

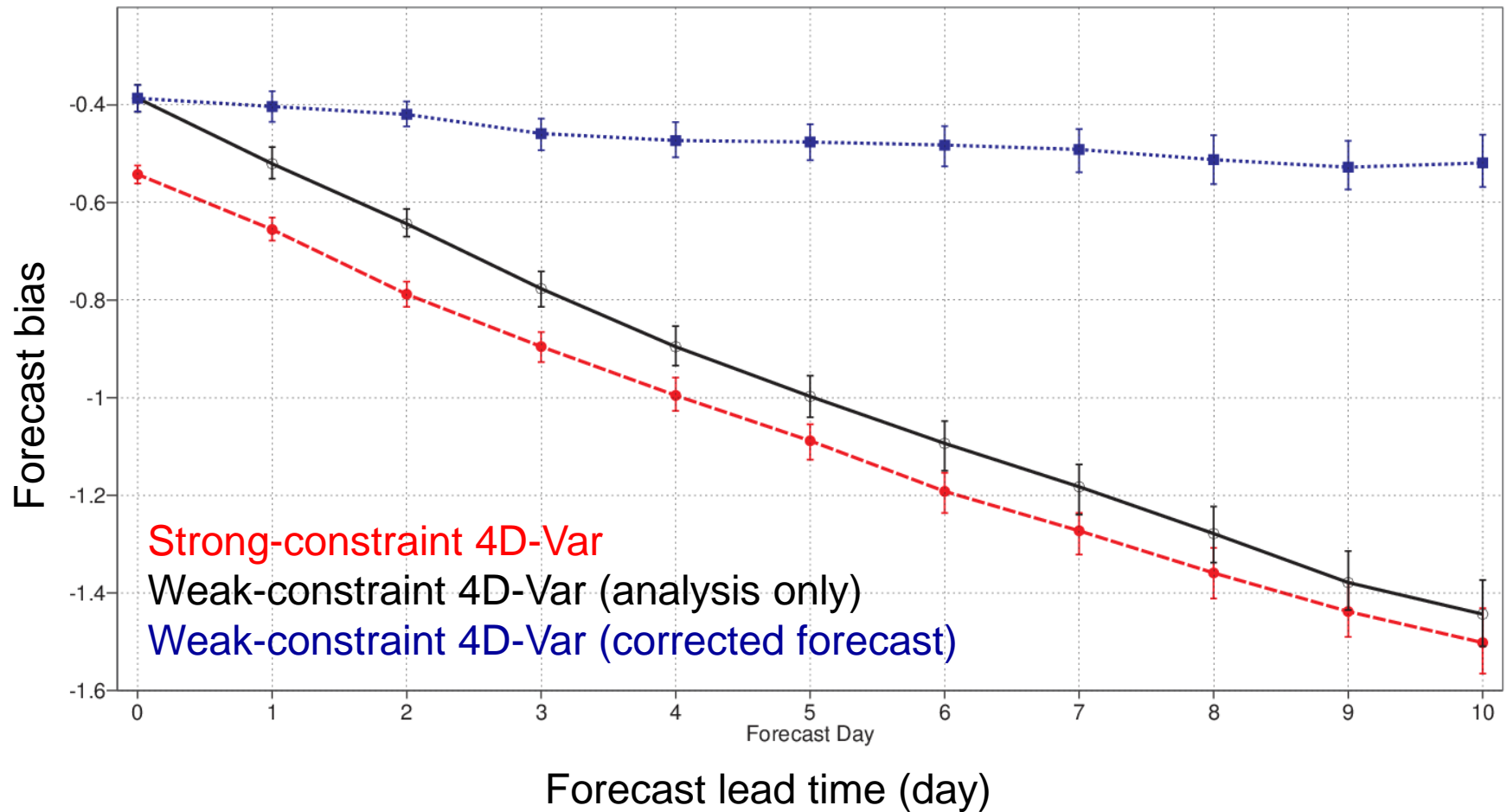
Mean forecast error against radiosonde temperature (70hPa)



The better analysis can be seen in the forecasts

Impact of new weak-constraint 4D-Var forcing on forecast skills

Mean forecast error against radiosonde temperature (70hPa)



- The model error estimation η is applied as a constant model forcing over 10 days
- The forecast model is not biased anymore and mean error does not increase
- Approach will be tested in seasonal forecast (some model error variability required)

Conclusions and future work

New application of RO at ECMWF

- Diagnose model deficiencies and correct them in 4D-Var
- Bias reduced up to 50% with weak-constraint 4D-Var (from -0.6C to -0.3C)

How many RO do we need to estimate the model bias? Homogeneous network?

What is the impact of weak-constraint 4D-Var on the observation bias correction?

What is the best way to correct forecast/reanalysis bias? Climatology for η ?